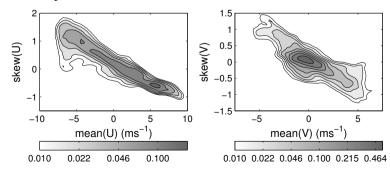
Case Study: Sea Surface Winds

- Air/Sea Exchange
 - ocean and atmosphere interact through respective boundary layers, exchanging momentum, energy, freshwater, and gases
 - fluxes depend on surface winds, in general nonlinearly
 - ocean currents largely driven by surface winds
- Sea State
 - sea state important for shipping, recreation
 - determined by both local and remote winds
- Power Generation
 - wind power potentially significant source of energy
 - generation rate scales as cube of wind speed; extreme events important

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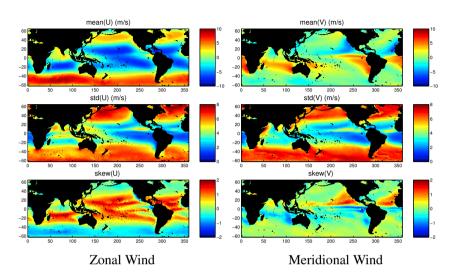
Mean and Skewness of Vector Wind

■ Joint pdfs of mean and skew for zonal and meridional winds



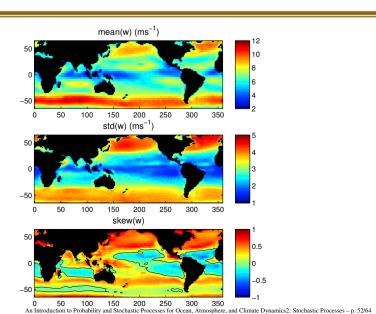
(note logarithmic contour scale)

Vector Wind Moments



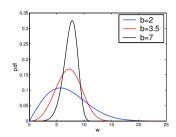
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Wind Speed Moments



Wind Speed pdf: Weibull distribution

 \blacksquare The pdf of wind speed w has traditionally (and empirically) been represented by 2-parameter Weibull distribution:



$$p(w) = \frac{b}{a} \left(\frac{w}{a}\right)^{b-1} \exp\left[-\left(\frac{w}{a}\right)^{b}\right]$$

- \blacksquare a is the <u>scale</u> parameter (pdf centre)
- b is the shape parameter (pdf tilt)
- $p_w(w)$ is unimodal

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Boundary Layer Dynamics

■ Horizontal momentum equations:

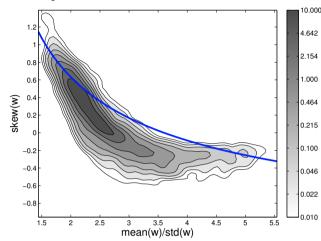
$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} - \frac{1}{\rho} \frac{\partial (\rho \overline{\mathbf{u}' u_3'})}{\partial z}$$

- Momentum tendency due to:
 - advection (transport by flow; secondary importance on daily timescales)
 - pressure gradient force
 - Coriolis force
 - turbulent momentum flux (in vertical)
- Integrated momentum budget over boundary layer *h*:

$$\frac{d\mathbf{u}}{dt} = -\frac{1}{\rho}\nabla p - f\hat{\mathbf{k}} \times \mathbf{u} + \frac{1}{h}\left(\overline{\mathbf{u}'u_3'}(0) - \overline{\mathbf{u}'u_3'}(h)\right)$$

Wind Speed pdfs: Observed

■ Observed speed moments fall around Weibull curve



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Surface Wind Stress

■ Surface wind stress is turbulent momentum flux across air/sea interface:

$$\tau_s = \rho_a \overline{\mathbf{u}' u_3'}(0)$$

along-mean wind component

cross-mean wind component

where (u,v)

= vertical wind component

Flux parameterised in terms of **u** by bulk drag formula:

$$\tau_s = \rho_a c_d w \mathbf{u}$$

where $w = ||\mathbf{u}||$ is the wind speed.

Surface Momentum Budget

- To close momentum budget, need parameterisation of turbulent momentum flux at z = h
- Use specified "entrainment velocity" W_e

$$\overline{\mathbf{u}'u_3'}(h) = W_e(\mathbf{U} - \mathbf{u})$$

⇒ Surface layer momentum budget

$$\frac{d\mathbf{u}}{dt} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} - \frac{c_d}{h} w \mathbf{u} + \frac{W_e}{h} (\mathbf{U} - \mathbf{u})$$

$$= \mathbf{\Pi} - \frac{c_d}{h} w \mathbf{u} - \frac{W_e}{h} \mathbf{u}$$

where

$$\mathbf{\Pi} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} + \frac{W_e}{h} \mathbf{U}$$

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Mechanistic Model: pdf

■ Solution of associated Fokker-Planck equation for stationary pdf:

$$p_{uv}(u,v) = \mathcal{N}_1 \exp\left(\frac{2}{\sigma^2} \left\{ \langle \Pi_u \rangle u - \frac{W_e}{2h} (u^2 + v^2) - \frac{1}{h} \int_0^{\sqrt{u^2 + v^2}} c_d(w') w'^2 dw' \right\} \right)$$

Changing to polar coordinates and integrating over angle gives wind speed pdf:

$$p_w(w) = \mathcal{N}wI_0\left(\frac{2\langle \Pi_u\rangle w}{\sigma^2}\right) \exp\left(-\frac{2}{\sigma^2}\left\{\frac{W_e}{2h}w^2 + \frac{1}{h}\int_0^w c_d(w')w'^2 dw'\right\}\right)$$

Mechanistic Model: SDE

■ Decomposing Π into mean and fluctuations:

$$\Pi_u(t) = \langle \Pi_u \rangle + \sigma \dot{W}_1(t)$$

$$\Pi_v(t) = \sigma \dot{W}_2(t)$$

where \dot{W}_i is Gaussian white noise

$$\left\langle \dot{W}_i(t_1)\dot{W}_j(t_2)\right\rangle = \delta_{ij}\delta(t_1 - t_2)$$

we obtain stochastic differential equation

$$\frac{du}{dt} = \langle \Pi_u \rangle - \frac{c_d}{h} wu - \frac{W_e}{h} u + \sigma \dot{W}_1$$

$$\frac{dv}{dt} = -\frac{c_d}{h} wv - \frac{W_e}{h} v + \sigma \dot{W}_2$$

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Mechanistic Model: Predictions

