

## SDEs: A first example

- We see in this example
  - i. a deep connection between diffusion of probability described by FPE and pathwise evolution described by SDE
  - ii. a connection between the “drift” and the “deterministic” part of the dynamics
  - iii. a connection between the “diffusion” and the “stochastic” part of the dynamics
  - iv. the stationary pdf includes contributions from both deterministic & stochastic terms. Fluctuations drive system away from deterministic attractor, dynamics pushes it back: stationary pdf balances these tendencies.
- These ideas important enough to help get Albert Einstein his Nobel Prize (for early work on the subject) ...

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 25/64

## Frankignoul & Hasselmann 1977

- These ideas also at heart of first **stochastic climate model** (Frankignoul & Hasselmann 1977), used to explain why SST spectra are generally red in character
- Simple model: SST anomalies driven by “fast” atmospheric forcing (assumed to be white in time) and “slow” relaxation to climatology through ocean mixed layer physics
- With  $T =$  SST anomaly, have model

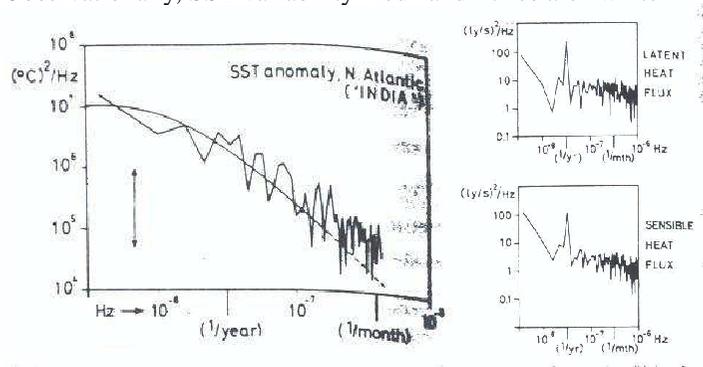
$$\frac{dT}{dt} = -\lambda T + \gamma \dot{W}$$

which predicts red SST response to fast, rapidly-decorrelating atmospheric forcing

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 26/64

## Frankignoul & Hasselmann 1977

- Observationally, SST variability “red” and fluxes are “white”



From Frankignoul & Hasselmann *Tellus* 1977

- Linear stochastic model  $\Rightarrow$  simple null hypothesis for observed variability

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 27/64

## Multivariate Ornstein-Uhlenbeck processes

- Have so far focused on scalar stochastic processes; can generalise results so far for  $\mathbf{x} \in \mathbb{R}^N$
- Vector white noise  $\dot{\mathbf{W}}(t)$  is an array of independent white noise processes

$$\mathbb{E}\{\dot{W}_i(t)\dot{W}_j(t')\} = \delta_{ij}\delta(t-t')$$

- For constant matrices  $A$  and  $B$ , SDE

$$\frac{d}{dt}\mathbf{x} = A\mathbf{x} + B\dot{\mathbf{W}}$$

has solution (with  $P(t) = e^{At}$ )

$$\mathbf{x}(t) = P(t)\mathbf{x}(0) + \int_0^t P(t-t')B\dot{\mathbf{W}}(t') dt'$$

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 28/64

## Multivariate Ornstein-Uhlenbeck processes

- Solution is Gaussian with mean and stationary autocovariance

$$\begin{aligned} E\{\mathbf{x}(t)\} &= P(t)\mathbf{x}(0) \\ C_{xx}(\tau) &= E\{\mathbf{x}(t+\tau)\mathbf{x}^T(t)\} = P(\tau)C_{xx}(0) \end{aligned}$$

where stationary covariance satisfies **Lyapunov equation**

$$AC_{xx}(0) + C_{xx}(0)A^T = -BB^T$$

⇒ balance between deterministic dynamics ( $A$ ), stochastic forcing ( $B$ ) and statistics of response ( $C_{xx}(0)$ ); an example of a **fluctuation-dissipation relationship**

- pdf  $p(\mathbf{x})$  satisfies Fokker-Planck equation

$$\partial_t p = - \sum_i A_i \partial_i p + \sum_{ij} (BB^T)_{ij} \partial_{ij}^2 p$$

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 29/64

## Optimal Perturbations

- If dynamical matrix  $A$  is *non-normal*, i.e.

$$AA^T \neq A^T A$$

so eigenvectors of  $A$  are not orthogonal, then:

- EOFs (eigenvectors of covariance  $C_{xx}(0)$ ) do not coincide with eigenvectors of  $A$  (dynamical modes)
- perturbation norm

$$N(t) = \mathbf{x}(t)^T M \mathbf{x}(t) = \mathbf{x}(0)^T P(t)^T M P(t) \mathbf{x}(0)$$

may grow (by potentially large amount) over finite times even though asymptotically stable:

$$\lim_{t \rightarrow \infty} N(t) = 0$$

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 30/64

## Optimal Perturbations

- Defining amplification factor

$$n(t) = \frac{\mathbf{x}(0)^T P(t)^T M P(t) \mathbf{x}(0)}{\mathbf{x}(0)^T M \mathbf{x}(0)}$$

optimal perturbation  $\mathbf{e}$  maximises  $n(t)$  subject to constraint  $\mathbf{x}(0)^T M \mathbf{x}(0) = 1 \Rightarrow$  generalised eigenvalue problem:

$$P(t)^T M P(t) \mathbf{e} = \lambda M \mathbf{e}$$

- Response to fluctuating forcing:

$$\text{var}(\mathbf{x}^T(t) M \mathbf{x}(t)) = B^T \left( \int_0^t P(s)^T M P(s) ds \right) B$$

⇒ importance of projection of noise structure on “average” optimals for maintaining variance; “stochastic optimals”

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 31/64

## Linear inverse modelling

- How are  $A$  and  $B$  determined in practice?

### 1. Empirical: Linear Inverse Modelling

- estimate covariances from observations and compute

$$A = \frac{1}{\tau} \ln (C_{xx}(\tau) C_{xx}(0)^{-1})$$

- Compute  $B$  from Lyapunov equation

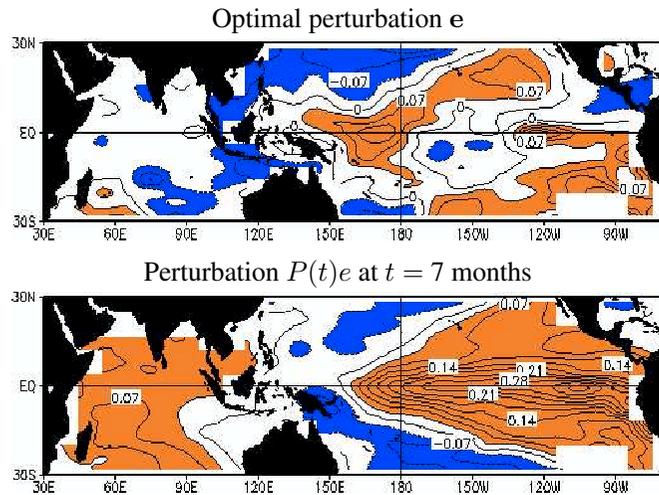
$$AC_{xx}(0) + C_{xx}(0)A^T = -BB^T$$

- Issues:

- if  $\mathbf{x}(t)$  not truly Markov, estimates will depend on lag  $\tau$
- must enforce positive-definiteness of  $BB^T$

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 32/64

## LIM: ENSO SST Optimal Perturbations



From Penland and Sardeshmukh (1995)

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 33/64

## Stochastic ENSO Models

### 2. Mechanistic: Stochastic Reduction of Tangent Linear Model

- Partition dynamics into “slow” and “fast” variables  $\mathbf{x}$  and  $\mathbf{y}$ :

$$\begin{aligned}\frac{d\mathbf{x}}{dt} &= \mathbf{f}(\mathbf{x}, \mathbf{y}) \\ \frac{d\mathbf{y}}{dt} &= \frac{1}{\epsilon} \mathbf{g}(\mathbf{x}, \mathbf{y})\end{aligned}$$

- Reduce coupled system to effective stochastic dynamics for  $\mathbf{x}$ :

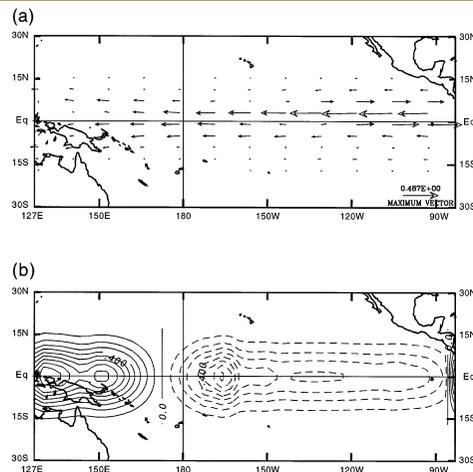
$$\frac{d\mathbf{x}}{dt} = L\mathbf{x} + N(\mathbf{x}, \mathbf{x}) + S(\mathbf{x}) \circ \dot{\mathbf{W}}$$

(many ways of doing this; some formal, some ad hoc)

- Linearise model around appropriate state (e.g. climatological mean)

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 34/64

## Mechanistic Model: 6-Month Stochastic Optimals



From Kleeman and Moore, *J. Atmos. Sci.*, 1997.

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 35/64

## SDEs: Multiplicative vs. additive noise

- General 1d SDE

$$\frac{dx}{dt} = a(x, t) + b(x, t)\dot{W}$$

- Have so far focused on case of constant noise strength  $b(x, t)$  (that is, noise independent of state of system)
- In such a case, noise said to be **additive**
- In general, however,  $b(x, t)$  depends on  $x$  and noise is **multiplicative**
- Things become much more complicated ...
- We can formally integrate the SDE:

$$x(t) = x(0) + \int_0^t a(x, t)dt + \int_0^t b(x, t)dW(t)$$

but we need to make clear what is meant by stochastic integral

An Introduction to Probability and Stochastic Processes for Ocean, Atmosphere, and Climate Dynamics2: Stochastic Processes – p. 36/64