

# Reconciling Non-Gaussian Climate Statistics with Linear Dynamics

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Gaussian statistics are consistent with linear dynamics.

Are non-Gaussian statistics necessarily inconsistent with linear dynamics ?

In particular, are skewed pdfs, implying different behavior of positive and negative anomalies, inconsistent with linear dynamics ?

*Thanks also to : Barsugli, Compo, Newman, Penland, and Shin*

# The Linear Stochastically Forced (LSF) Approximation

$$\frac{dx}{dt} = A x + f_{ext} + B \eta$$

$x$  =  $N$ -component anomaly state vector  
 $\eta$  =  $M$ -component gaussian noise vector  
 $f_{ext}(t)$  =  $N$ -component external forcing vector  
 $A(t)$  =  $N \times N$  matrix  
 $B(t)$  =  $N \times M$  matrix

## Supporting Evidence

- Linearity of coupled GCM responses to radiative forcings
- Linearity of atmospheric GCM responses to tropical SST forcing
- Linear dynamics of observed seasonal tropical SST anomalies
- Competitiveness of linear seasonal forecast models with global coupled models
- Linear dynamics of observed weekly-averaged circulation anomalies
- Competitiveness of Week 2 and Week 3 linear forecast models with NWP models
- Ability to represent observed second-order synoptic-eddy statistics

# Observed and Simulated Spectra of Tropical SST Variability

Spectra of the projection of tropical SST anomaly fields on the 1st EOF of observed monthly SST variability in 1950-1999.

**Observations (Purple)**

**IPCC AR4 coupled GCMs**

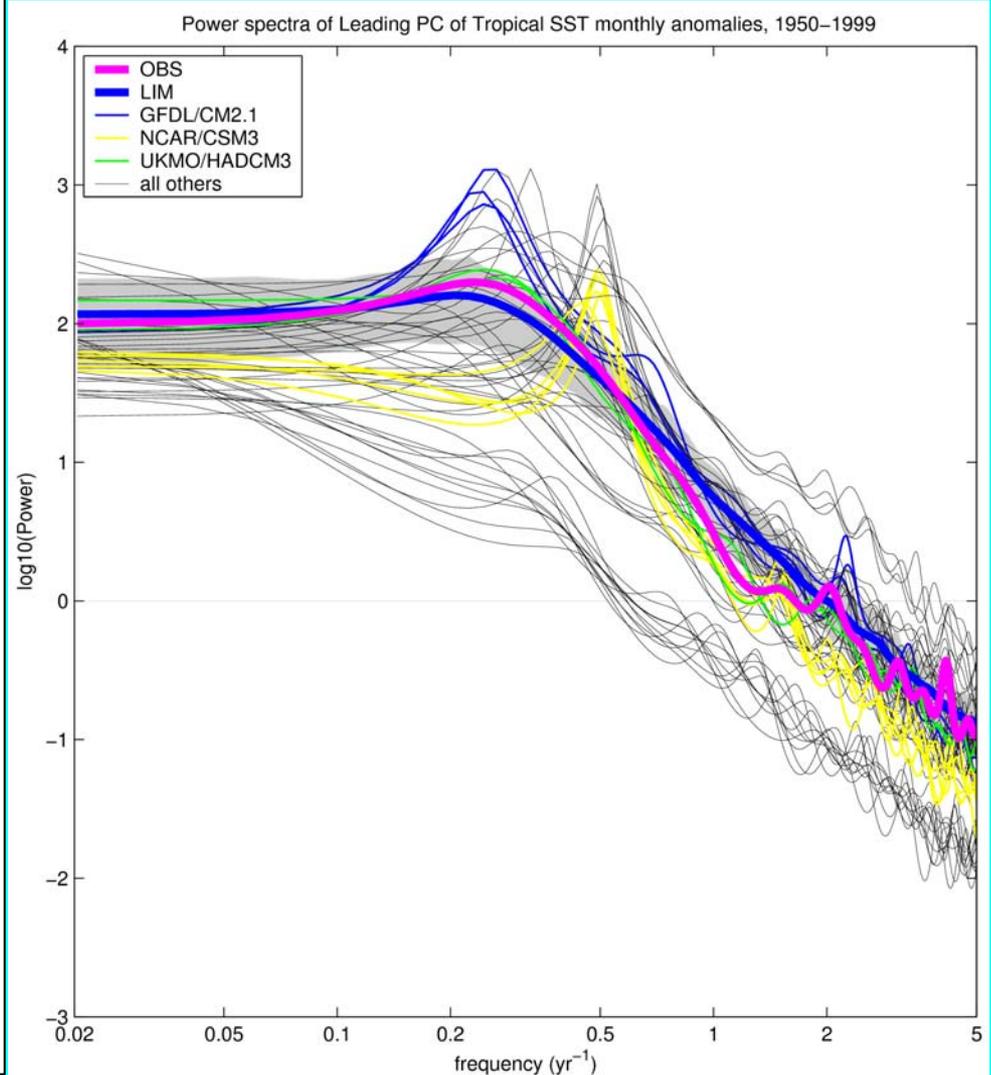
(20<sup>th</sup>-century (20c3m) runs)

(thin black, yellow, blue, and green)

**A linear inverse model (LIM)** constructed from 1-week lag covariances of weekly-averaged tropical data in 1982-2005 (Thick Blue)

Gray Shading :

95% confidence interval from the LIM, based on 100 model runs with different realizations of the stochastic forcing.



*From Newman, Sardeshmukh and Penland (2008)*

# Seasonal Predictions of Ocean Temperatures in the Eastern Tropical Pacific : Comparison of linear empirical and nonlinear GCM forecast skill

*(Courtesy : NCEP)*

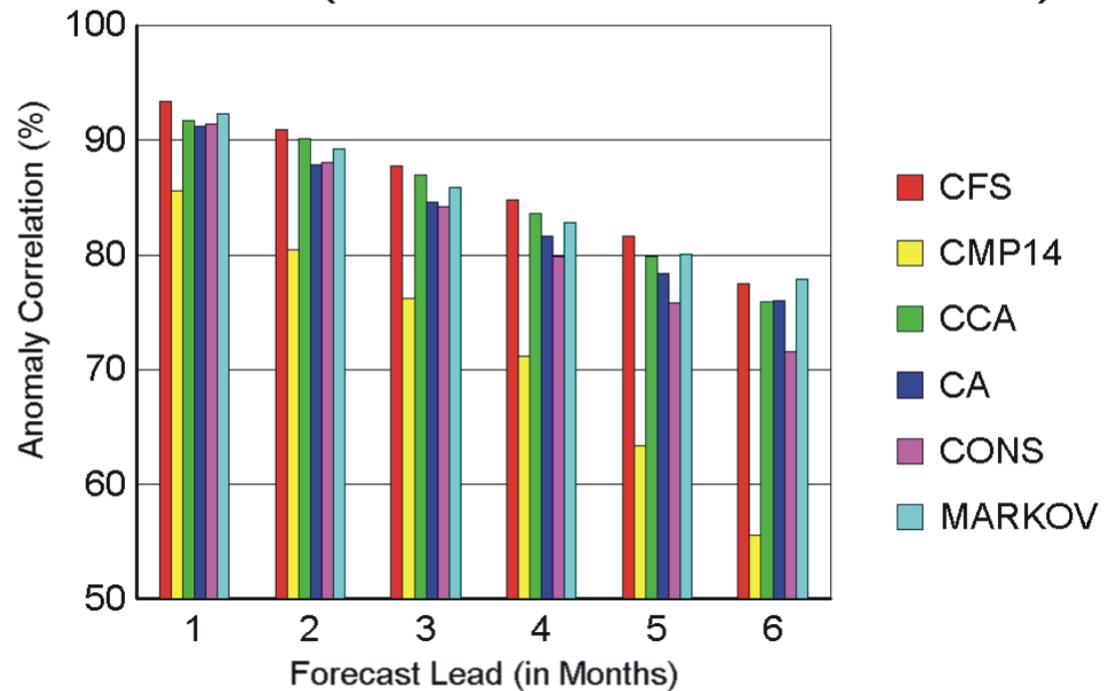
Simple linear  
empirical  
models are  
apparently  
just as good  
at predicting  
ENSO

as

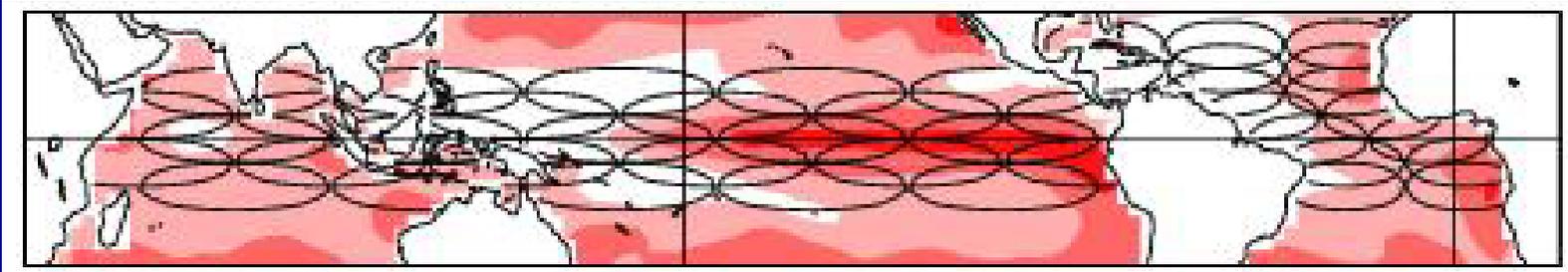
“state of  
the art”  
coupled

GCMs

## Skill in SST Anomaly Prediction Nino-3.4 (DJF 97/98 to DJF 03/04)

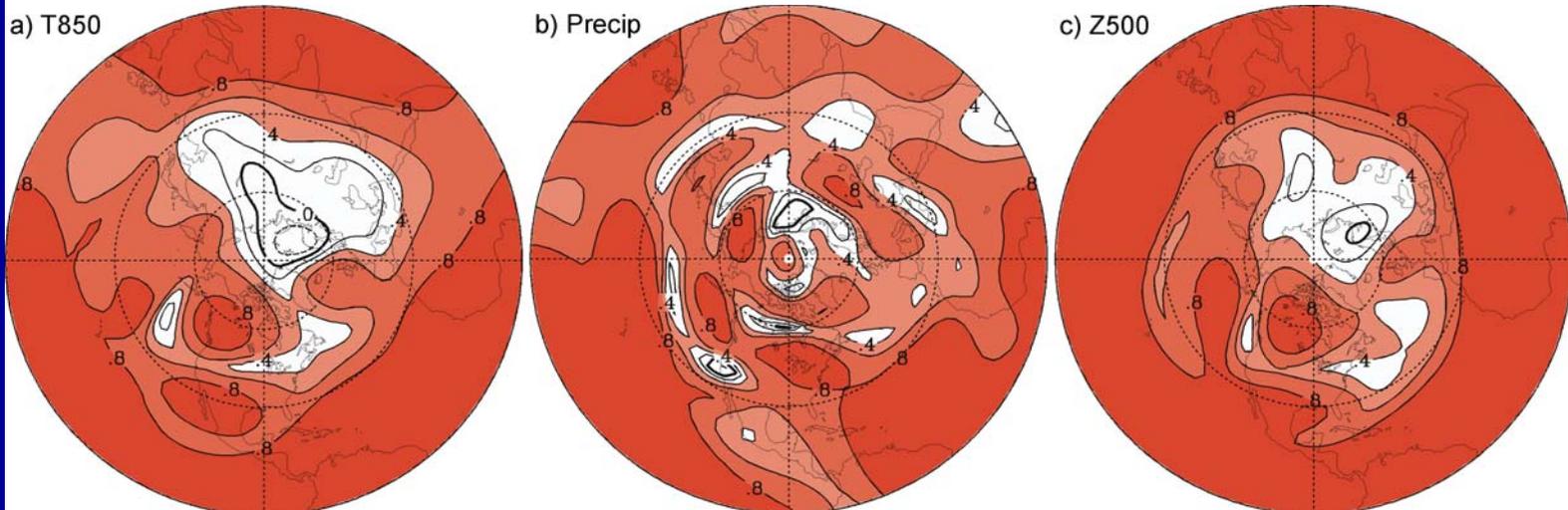


# DOMINANCE and LINEARITY of Tropical SST influences on global climate variability



**BASIC POINT:** The nonlinear NCAR/CCM3 atmospheric GCM's responses to prescribed global SST changes over the last 50 years are well -approximated by linear responses to just the Tropical SST changes, obtained by linearly combining the GCM's responses to SSTs in the 43 localized areas shown above.

## Local correlation of annual mean “GOGA” and “Linear TOGA” responses



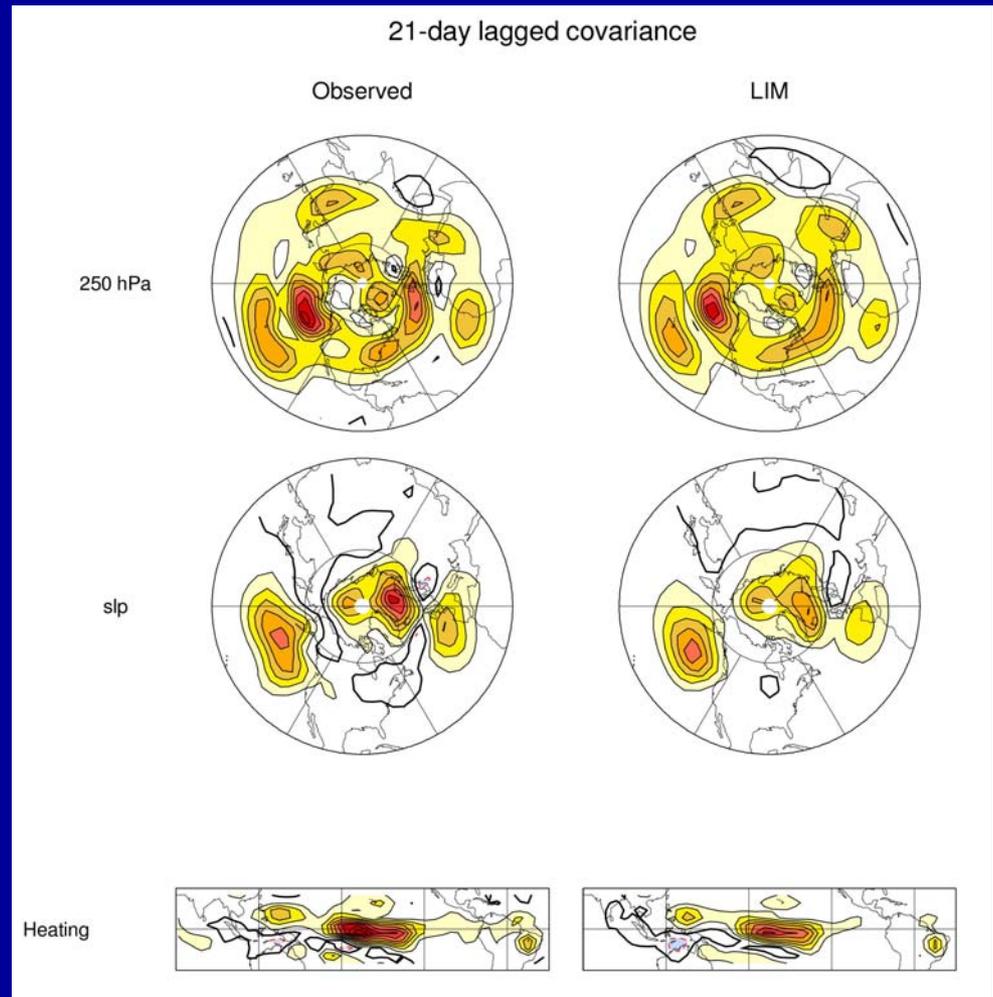
# Decay of lag-covariances of weekly anomalies is consistent with linear dynamics

Is  $C(\tau) = e^{M\tau} C(0)$  ?

$M$  is first estimated using the observed  $C(\tau = 5 \text{ days})$  and  $C(0)$  in this equation, and then used to "predict"  $C(\tau = 21 \text{ days})$

The components of the anomaly state vector  $\mathbf{x}$  include the 7-day running mean PCs of 250 and 750 mb streamfunction, SLP, tropical diabatic heating and stratospheric height anomalies.

From Newman and Sardeshmukh (2008)



# An attractive feature of the LSF Approximation

$$\frac{dx}{dt} = A x + f_{ext} + B \eta$$

## Equations for the first two moments

(Applicable to both Marginal and Conditional Moments)

$\langle x \rangle$  = ensemble mean anomaly

$C$  = covariance of departures from ensemble mean

$$\frac{d}{dt} \langle x \rangle = A \langle x \rangle + f_{ext}$$

$$\frac{d}{dt} C = A C + C A^T + B B^T$$

If  $A(t)$ ,  $B(t)$ , and  $f_{ext}(t)$  are constant, then

First two **Marginal** moments



$$\langle x \rangle = -A^{-1} f_{ext}$$

$$\frac{dC}{dt} = 0 = A C + C A^T + B B^T$$

First two **Conditional** moments

Ensemble mean forecast

Ensemble spread



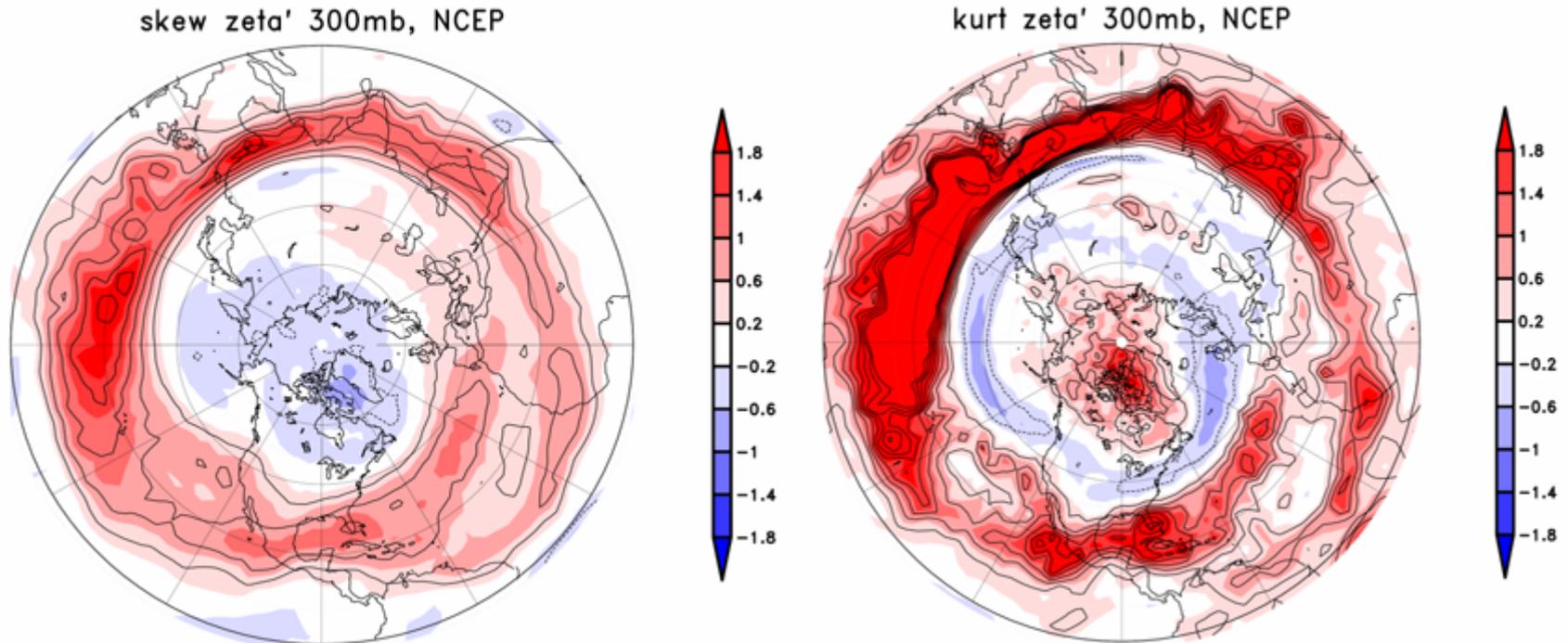
$$\hat{x}'(t) \equiv \langle x'(t) \mid x'(0) \rangle = e^{At} x'(0)$$

$$\hat{C}(t) \equiv \langle (\hat{x}' - x') (\hat{x}' - x')^T \rangle = C - e^{At} C e^{A^T t}$$

If  $x$  is Gaussian, then these moment equations **COMPLETELY** characterize system variability *and* predictability

*But . . . atmospheric circulation statistics are not Gaussian . . .*

**Observed Skew  $S$  and (excess) Kurtosis  $K$  of daily 300 mb Vorticity (DJF)**

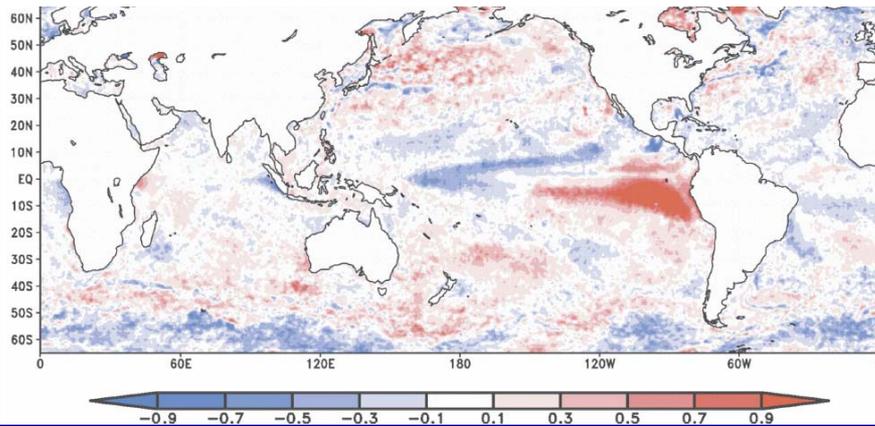


*From Sardeshmukh and Sura 2008*

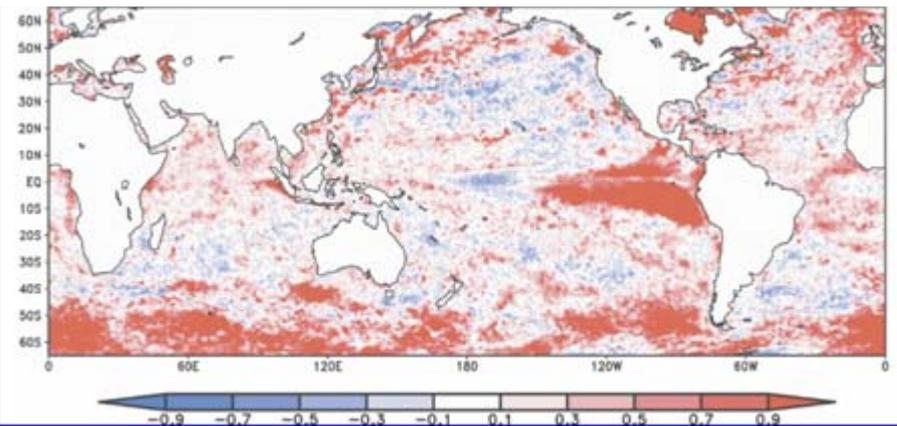
# *Sea Surface Temperature statistics are also not Gaussian . . .*

## **Observed Skew $S$ and (excess) Kurtosis $K$ of daily SSTs (DJF)**

**Skew**



**Kurtosis**



*From Sura and Sardeshmukh 2008*

# Modified LSF Dynamics

$$\text{Model 1: } \frac{dx}{dt} = Ax + f_{ext} + B\eta$$

$$\text{Model 2: } \frac{dx}{dt} = Ax + f_{ext} + B\eta + (Ex)\xi$$

$$\text{Model 3: } \frac{dx}{dt} = Ax + f_{ext} + B\eta + (Ex + g)\xi - \frac{1}{2}Eg$$

For simplicity consider a scalar  $\xi$  here

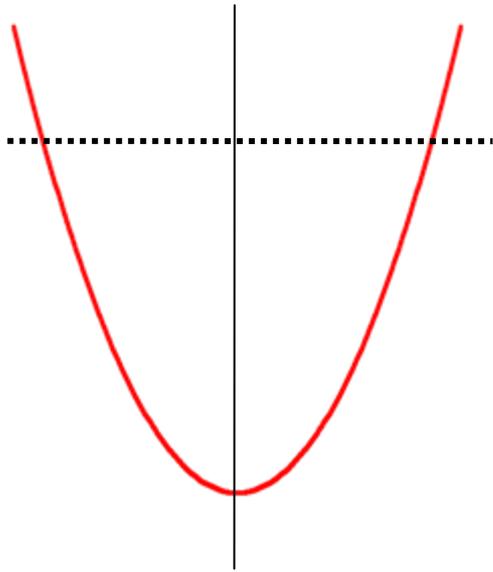
$A(t), B(t), E(t)$  are matrices;  $g(t), f_{ext}(t), \eta$  are vectors

*Moment Equations :*

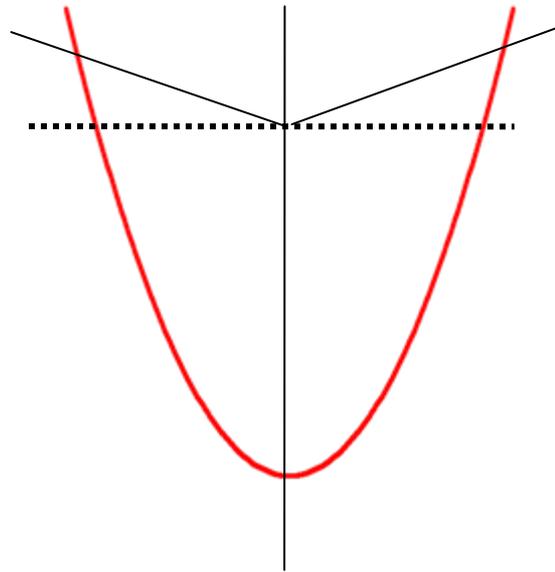
$$\frac{d}{dt} \langle x \rangle = M \langle x \rangle + f_{ext} \quad \text{where} \quad M = \left( A + \frac{1}{2} E^2 \right)$$

$$\frac{d}{dt} C = M C + C M^T + B B^T + E \{ C + \langle x \rangle \langle x \rangle^T \} E^T + g g^T$$

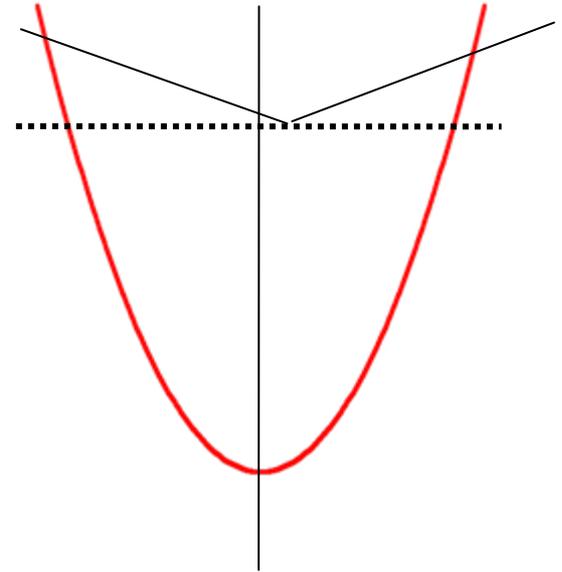
**A simple view of how additive and linear multiplicative noise can generate skewed PDFs even in a deterministically linear system**



**Additive noise only**  
**Gaussian**  
**No skew**



**Additive and uncorrelated**  
**Multiplicative noise**  
**Symmetric non-Gaussian**



**Additive and correlated**  
**Multiplicative noise**  
**Asymmetric non-Gaussian**

## A simple rationale for Correlated Additive and Multiplicative (CAM) noise

In a quadratically nonlinear system with “slow” and “fast” components  $x$  and  $y$ , the anomalous nonlinear tendency has terms of the form :

$$\begin{aligned}
 (xy)' &= x' \bar{y} + \bar{x} y' + x' y' - \overline{x' y'} \\
 &= \bar{y} x' + (\bar{x} + x') y' - \overline{x' y'}
 \end{aligned}$$

**CAM  
noise**

**mean Noise  
Induced Drift**

Note that it is the *STOCHASTICITY* of  $y'$  that enables the mean drift to be parameterized in terms of the noise amplitude parameters

$$\begin{aligned}
 (vT)' &= T' \bar{v} + v' \bar{T} + v' T' - \overline{v' T'} \\
 &= \bar{v} T' + (\bar{T} + T') v' - \overline{v' T'}
 \end{aligned}$$

# Rationalizing linear anomaly dynamics

## with correlated additive and linear multiplicative stochastic noise

$$\frac{dX_i}{dt} = L_{ij}X_j + N_{ijk}X_jX_k + F_i \quad \text{Einstein Summation Convention}$$

$$\frac{dX'_i}{dt} = [L_{ij} + (N_{ijk} + N_{ikj})\bar{X}_k] X'_j + N_{ijk}(X'_jX'_k - \overline{X'_jX'_k}) + F'_i$$

$$\text{Let } X' = \begin{bmatrix} x' \\ \eta' \end{bmatrix} \text{ and } \bar{X} = \begin{bmatrix} \bar{x} \\ \bar{\eta} \end{bmatrix}$$

$$\begin{aligned} \frac{dx'_i}{dt} = & [L_{ij} + (N_{ijm} + N_{imj})\bar{\eta}_m] x'_j && \text{Linear terms (= } A_{ij}x'_j) \\ & + [(N_{ijm} + N_{imj})x'_j + \{L_{im} + (N_{ijm} + N_{imj})\bar{x}_j\}] \eta'_m && \text{Correlated additive and multiplicative noise} \\ & - (N_{ijm} + N_{imj}) \overline{x'_j \eta'_m} && \text{Mean noise-induced drift} \\ & + N_{imn}(\eta'_m \eta'_n - \overline{\eta'_m \eta'_n}) && \text{Other additive noise (= } B_{ik} \xi_k) \\ & + N_{ijk} (x'_j x'_k - \overline{x'_j x'_k}) && \text{Hard nonlinearity} \\ & + f'_i && \text{External forcing} \end{aligned}$$

Neglecting the hard nonlinearity, and using the FPE to derive the noise-induced drift, we obtain

$$\begin{aligned} \frac{dx'_i}{dt} &= A_{ij} x'_j + (E_{ijm} x'_j + L_{im} + E_{ijm} \bar{x}_j) \eta'_m - \frac{1}{2} E_{ijm} (L_{jm} + E_{jkm} \bar{x}_k) + B_{ik} \xi_k + f'_i \\ &= A_{ij} x'_j + (E_{ijm} x'_j + G_{im}) \eta'_m - \frac{1}{2} E_{ijm} G_{jm} + B_{ik} \xi_k + f'_i \end{aligned}$$

where  $E_{ijm} = (N_{ijm} + N_{imj})$ , and  $G_{im} = L_{im} + (N_{ijm} + N_{imj})\bar{x}_j = L_{im} + E_{ijm}\bar{x}_j$

# A 1-D system with Correlated Additive and Multiplicative (“CAM”) noise

*Stochastic Differential Equation :*

$$\frac{dx}{dt} \cong Ax + (Ex + g)\eta + B\xi - \frac{1}{2}Eg$$

*Fokker-Planck Equation :*

$$Mxp = \frac{1}{2} \frac{d}{dx} [(E^2x^2 + 2Egx + g^2 + B^2) p]$$

*Moments :*

$$\langle x \rangle = 0$$

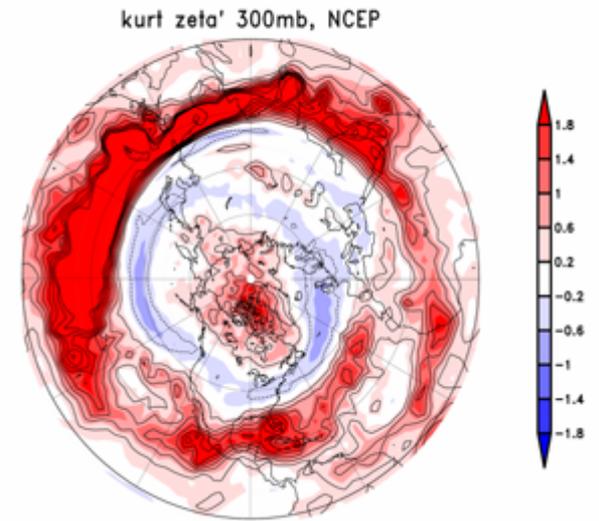
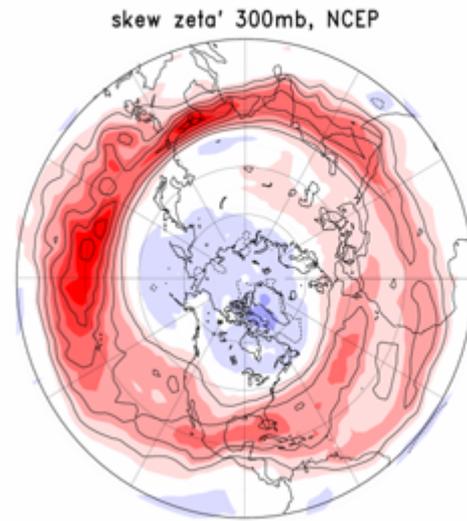
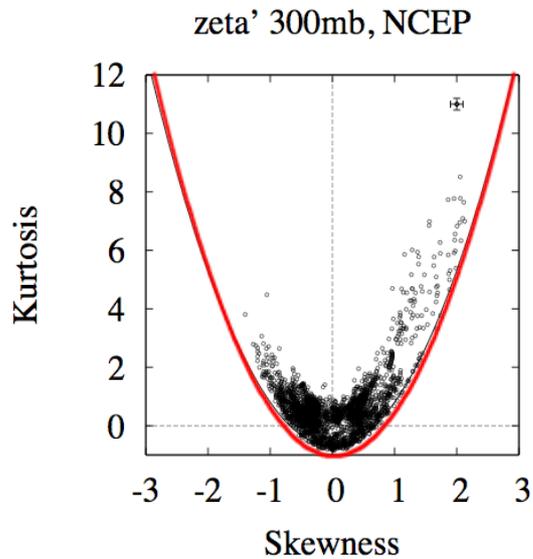
$$\langle x^n \rangle = - \left( \frac{n-1}{2} \right) [2Eg \langle x^{n-1} \rangle + (g^2 + B^2) \langle x^{n-2} \rangle] / \left[ M + \left( \frac{n-1}{2} \right) E^2 \right]$$

*A simple relationship between Skew and Kurtosis :*

Remembering that Skew  $S = \frac{\langle x^3 \rangle}{\sigma^3}$  and Kurtosis  $K = \frac{\langle x^4 \rangle}{\sigma^4} - 3$ , we have

$$K = \frac{3}{2} \left[ \frac{M + E^2}{M + (3/2)E^2} \right] S^2 + 3 \left[ \frac{M + (1/2)E^2}{M + (3/2)E^2} - 1 \right] \geq \frac{3}{2} S^2$$

## Observed Skew $S$ and (excess) Kurtosis $K$ of daily 300 mb Vorticity (DJF)

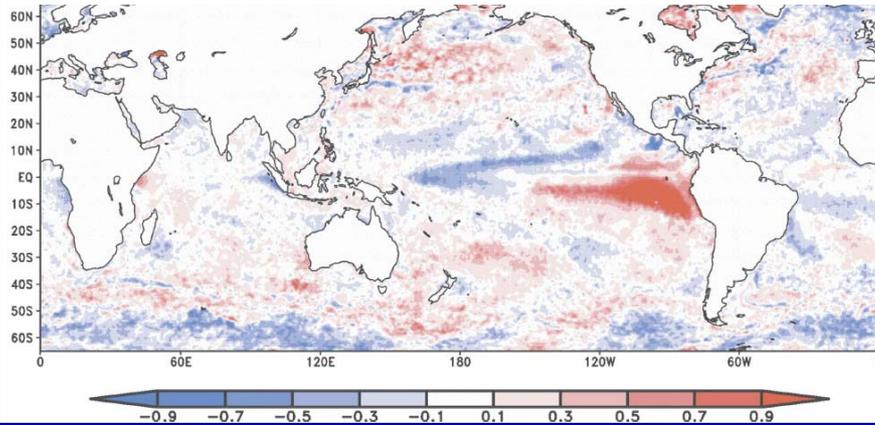


Note the quadratic relationship between  $K$  and  $S$  :

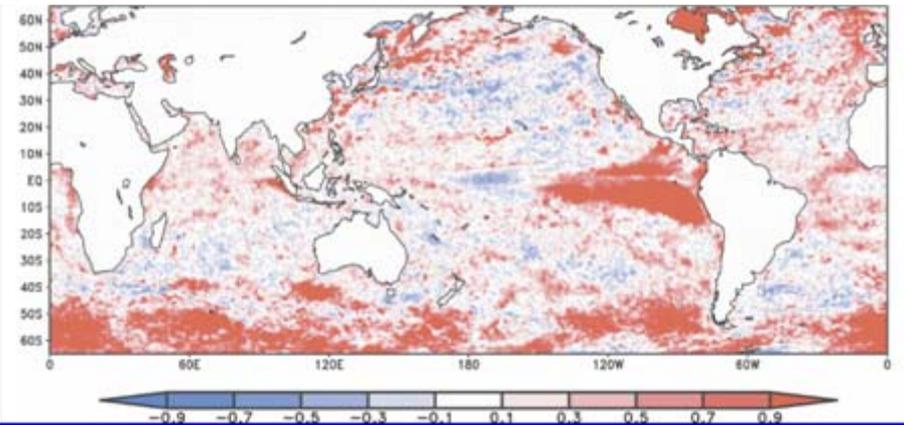
$$K \geq \frac{3}{2} S^2$$

# Observed Skew $S$ and (excess) Kurtosis $K$ of daily SSTs (DJF)

## Skew



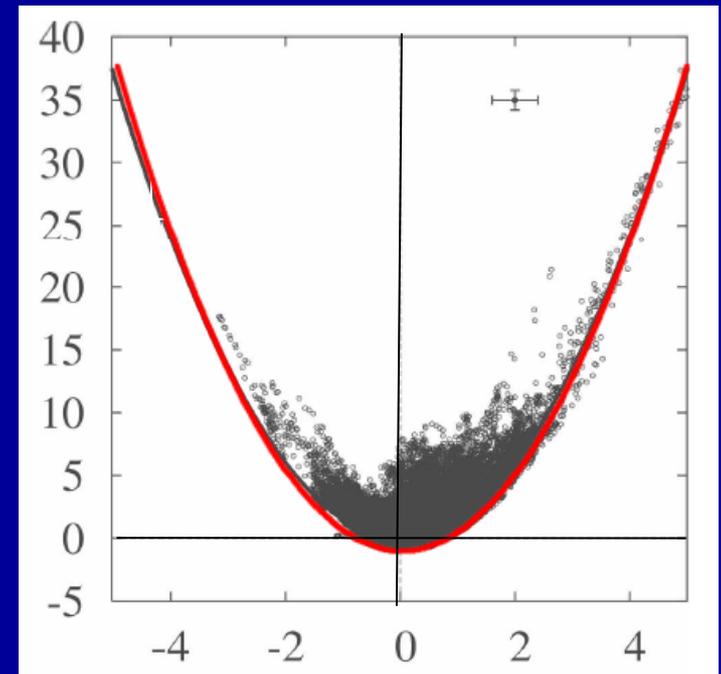
## Kurtosis



Note the quadratic relationship

between  $K$  and  $S$  :  $K \geq 3/2 S^2$

*From Sura and Sardeshmukh 2008*



# Understanding the patterns of Skewness and Kurtosis

**Are diabatic or adiabatic stochastic transients more important ?**

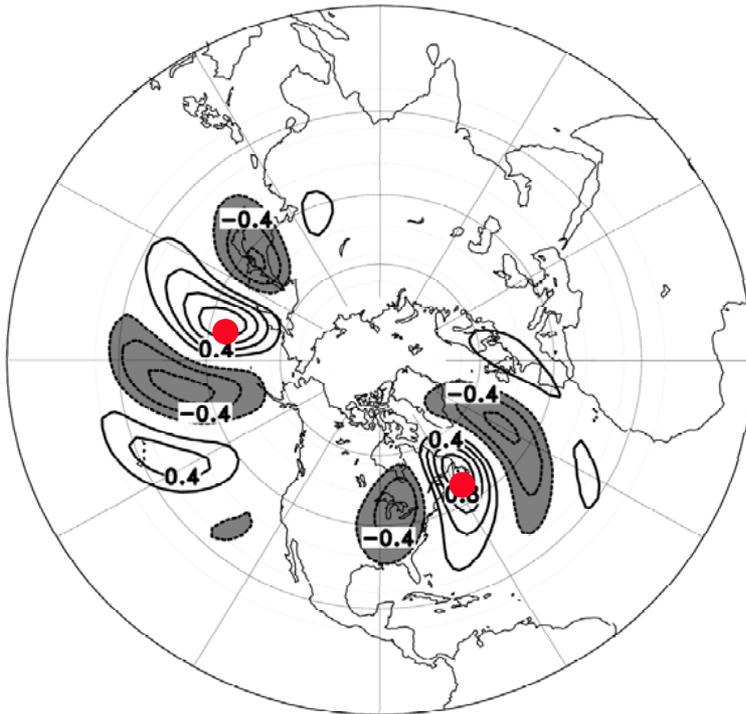
To clarify this, we examined the circulation statistics in a 1200 winter simulation generated with a T42 5-level dry adiabatic GCM (“PUMA”) with the observed time-mean diabatic forcing specified as a **fixed** forcing.

**There is thus NO transient diabatic forcing in these runs.**

# 1-point anomaly correlations of synoptic (2 to 6 day period) variations with respect to base points in the Pacific and Atlantic sectors

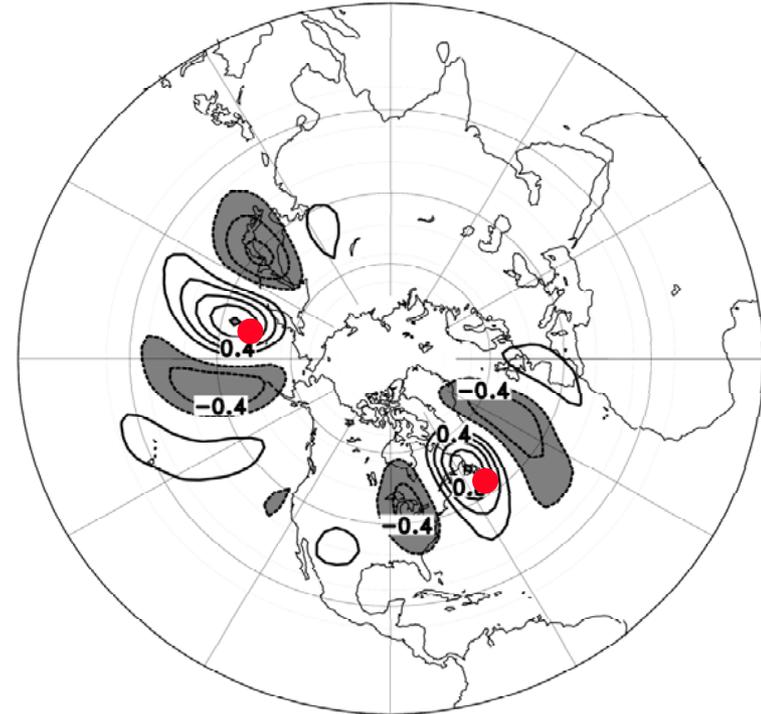
## Simulated

$z'$  500mb one-point correlations  
2-6 days, PUMA( $F_{\text{bar}}$ )

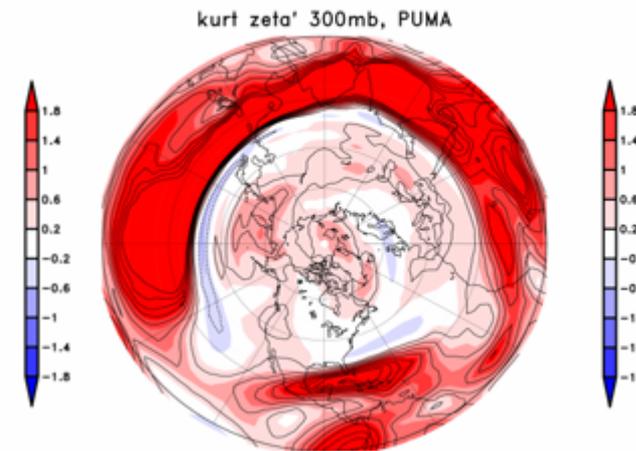
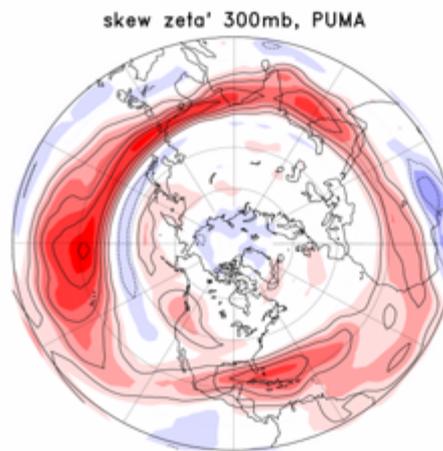
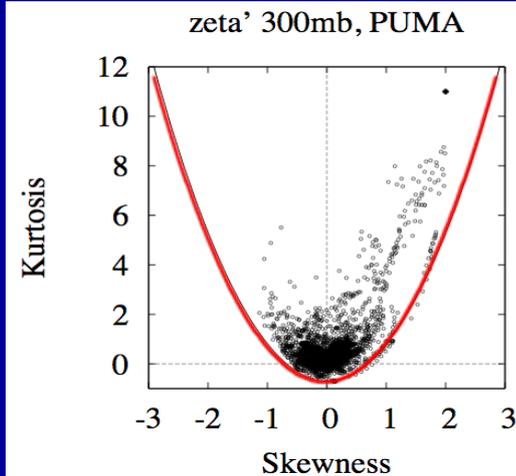
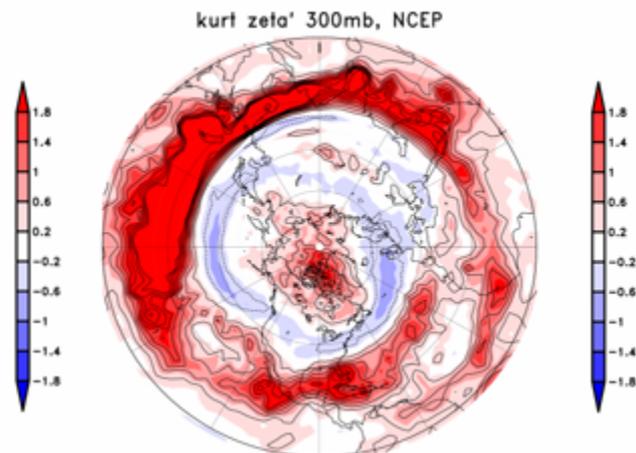
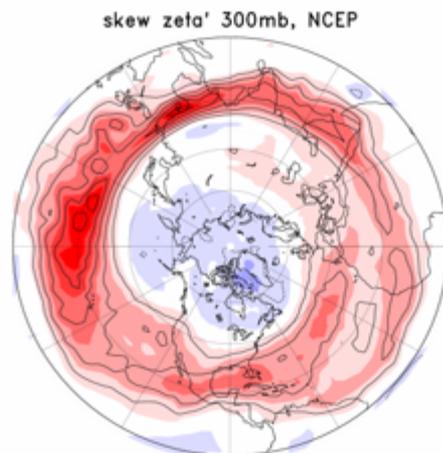
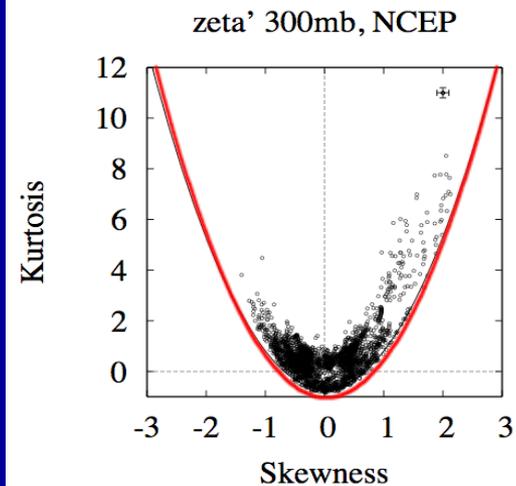


## Observed

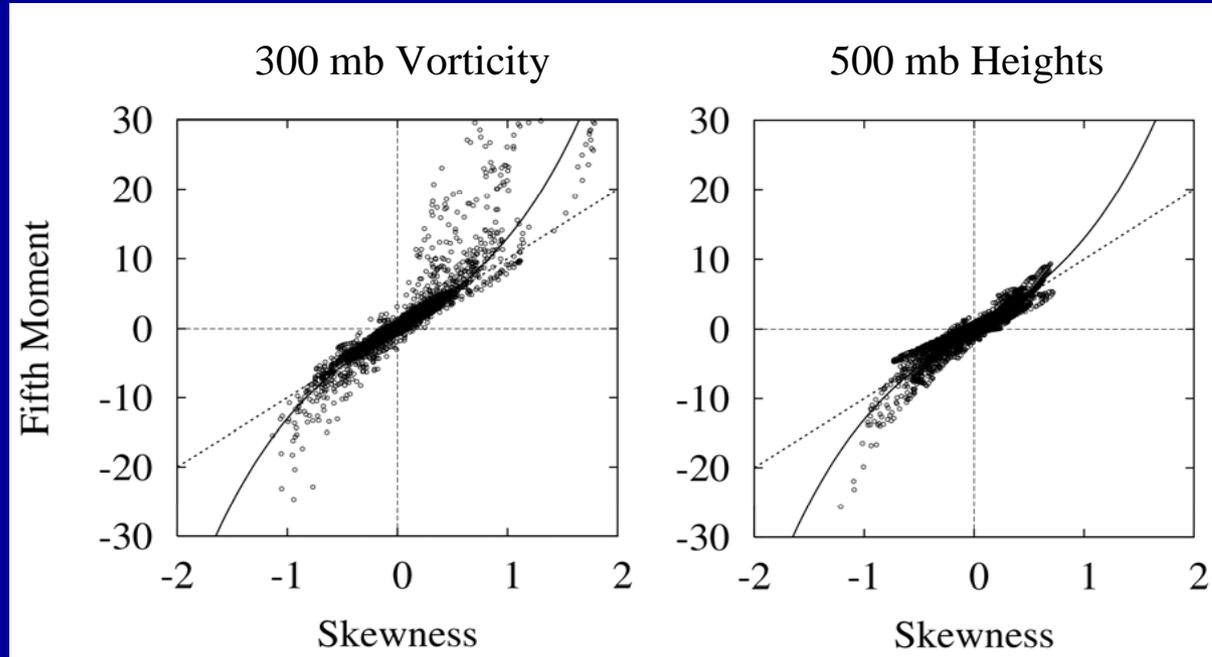
$z'$  500mb one-point correlations  
2-6 days, NCEP



# Observed (NCEP, Top) and Simulated (PUMA, Bottom) $S$ and $K$ of 300 mb Vorticity



# Scatter plots of **Fifth Moments versus Skew** in the dry adiabatic GCM



The 1-d model predicts  $\mu_5 \equiv \frac{\langle x^5 \rangle}{\sigma^5} > 10s + 3S^3$  for  $S > 0$   
<  $10s + 3S^3$  for  $S < 0$  !!

## A linear 1-D system with non-Gaussian statistics, forced by “CAM” noise

$$\frac{dx}{dt} = Ax + b\eta_1 + (Ex + g)\eta_2 - \frac{1}{2}Eg \quad \text{SDE}$$

$$[Mx]p = \frac{1}{2} \frac{d}{dx} [E^2x^2 + 2Egx + (g^2 + b^2)p] \quad \text{FPE}$$

$$p(x) = \frac{1}{N} \left[ (Ex + g)^2 + b^2 \right]^{\frac{1}{\alpha}-1} \exp \left[ -\frac{2g}{\alpha b} \arctan \left( \frac{Ex + g}{b} \right) \right] \quad \text{PDF}$$

Such a system satisfies  $K > (3/2)S^2$  and its PDF has power-law tails

$$M = A + 0.5 E^2$$

$$\alpha = E^2 / M$$

$$\text{Both} < 0$$

# Observed and Simulated pdfs in the North Pacific

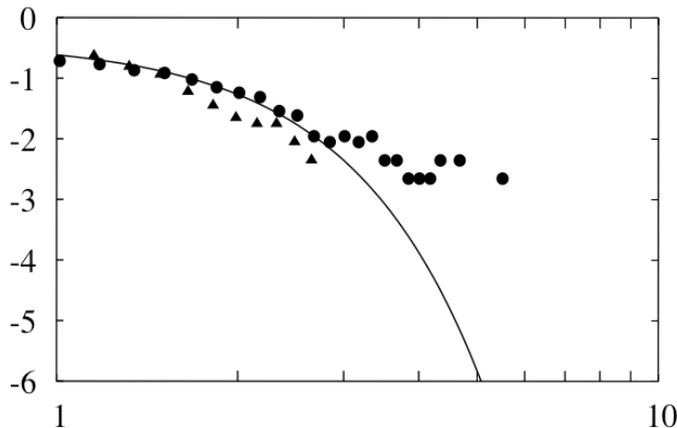
*(On a log-log plot, and with the negative half folded over into the positive half)*

**Observed  
(NCEP Reanalysis)**

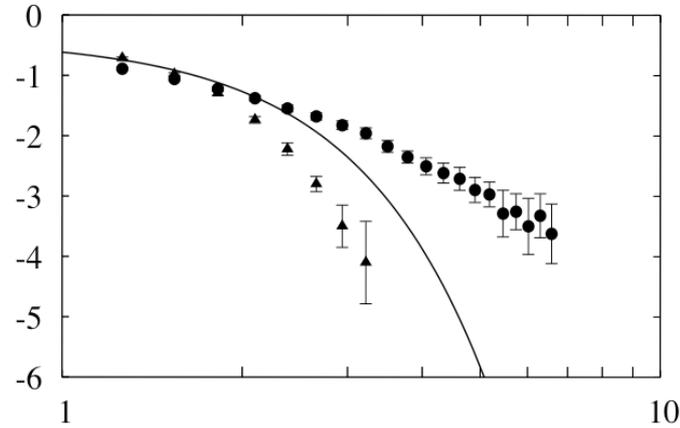
**Simulated by a dry adiabatic  
GCM with fixed forcing**

**500 mb  
Height**

z' 500mb, NCEP, 15N, 180W

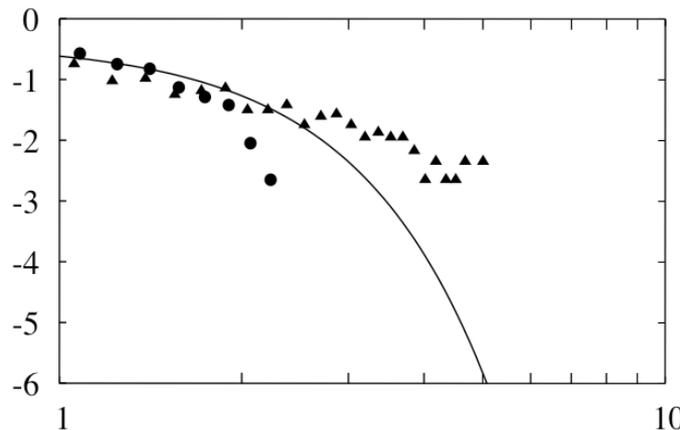


z' 500mb, PUMA, 15N, 180W

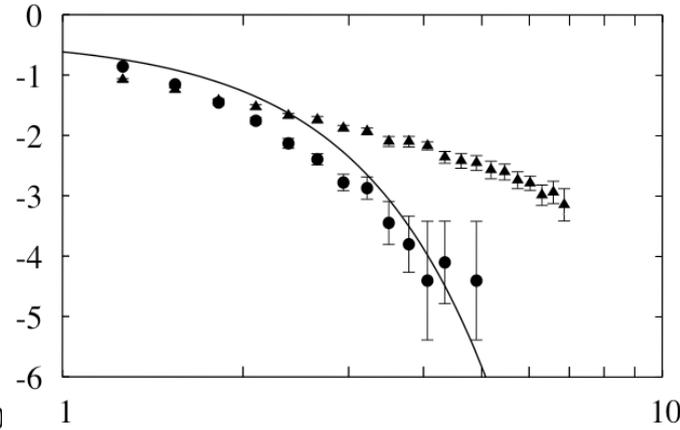


**300 mb  
Vorticity**

zeta' 300mb, NCEP, 20N, 180W



zeta' 300mb, PUMA, 20N, 180W



# Observed and Simulated pdfs in the North Pacific

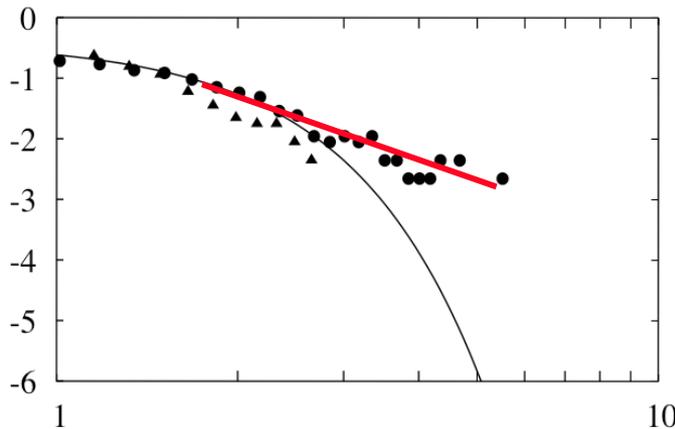
*(On a log-log plot, and with the negative half folded over into the positive half)*

**Observed  
(NCEP Reanalysis)**

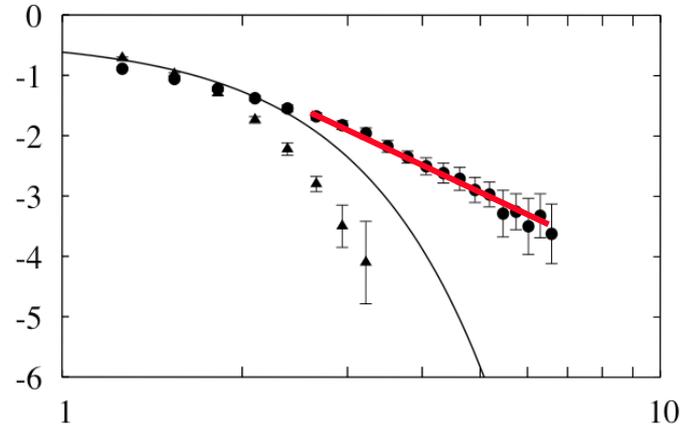
**Simulated by a dry adiabatic  
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**500 mb  
Height**

z' 500mb, NCEP, 15N, 180W

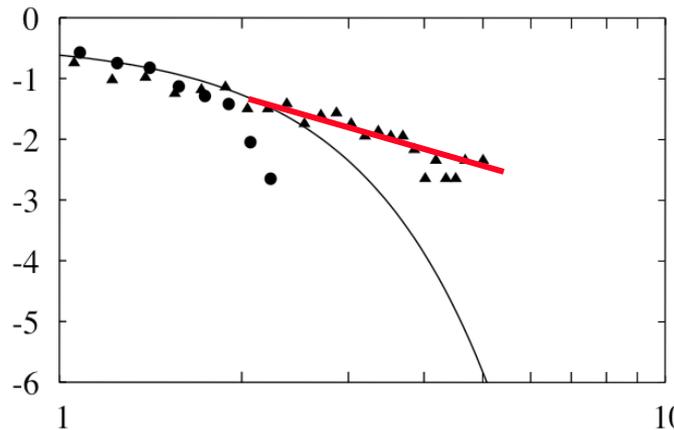


z' 500mb, PUMA, 15N, 180W

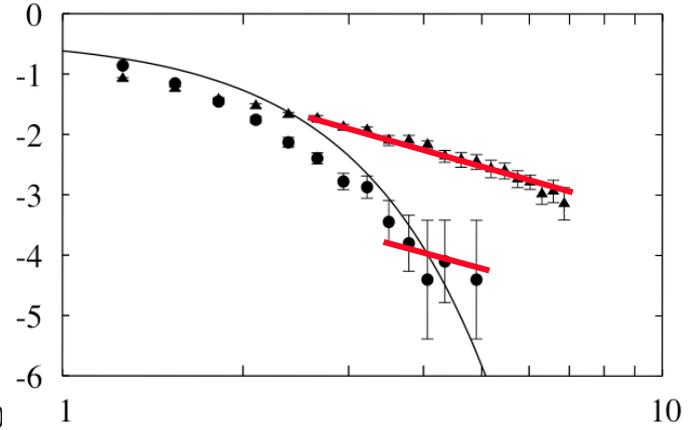


**300 mb  
Vorticity**

zeta' 300mb, NCEP, 20N, 180W



zeta' 300mb, PUMA, 20N, 180W



## A linear 1-D system with non-Gaussian statistics, forced by “CAM” noise

$$\frac{dx}{dt} = Ax + b\eta_1 + (Ex + g)\eta_2 - \frac{1}{2}Eg \quad \text{SDE}$$

$$[Mx]p = \frac{1}{2} \frac{d}{dx} [E^2x^2 + 2Egx + (g^2 + b^2)p] \quad \text{FPE}$$

$$p(x) = \frac{1}{N} \left[ (Ex + g)^2 + b^2 \right]^{\frac{1}{\alpha}-1} \exp \left[ -\frac{2g}{\alpha b} \arctan \left( \frac{Ex + g}{b} \right) \right] \quad \text{PDF}$$

Such a system satisfies  $K > (3/2)S^2$  and its PDF has power-law tails

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$$\text{Both} < 0$$

## A linear 1-D system with non-Gaussian statistics, forced by “CAM” noise

$$\frac{dx}{dt} = Ax + b\eta_1 + (Ex + g)\eta_2 - \frac{1}{2}Eg \quad \text{SDE}$$

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Such a system satisfies  $K > (3/2)S^2$  and its PDF has power-law tails

$$M = A + 0.5 E^2$$

$$\alpha = E^2 / M$$

$$\text{Both} < 0$$

## The most general linear 1-D system with non-Gaussian statistics, forced by “radical” noise

$$\frac{dx}{dt} = Ax + \sum_m \sqrt{[(E_m x + g_m)^2 + c_m x]} \eta_m - \frac{\beta}{2} + f_{ext} \quad \text{SDE}$$

$$[Mx + f_{ext}]p = \frac{1}{2} \frac{d}{dx} [(E^2 x^2 + 2\beta x + G^2)p] \quad \text{FPE}$$

$$p(x) = \frac{1}{N} [E^2 x^2 + 2\beta x + G^2]^{\frac{1}{\alpha}-1} \exp\left[\frac{2}{\gamma} \left(f_{ext} - \frac{\beta}{\alpha}\right) \arctan\left(\frac{E^2 x + \beta}{\gamma}\right)\right] \quad \text{PDF}$$

Such a system satisfies  $K \geq (3/2)S^2$  and its PDF also has power-law tails

$$\beta = \sum_m \left( E_m g_m + \frac{c_m}{2} \right)$$

$$E^2 = \sum_m E_m^2$$

$$G^2 = \sum_m g_m^2$$

**Why does a local 1-D system capture the relationships between the higher-order moments of the N-d climate system with obviously important non-local dynamics ?**

Mainly because the equations for the higher moments in the N-d system are increasingly dominated by **self-correlation** terms. We call this a principle of “**DIAGONAL DOMINANCE**”

$$K = \frac{3}{2} S^2 + r$$

$$r = 3 \left[ \frac{M + (1/2)E^2}{M + (3/2)E^2} - 1 \right] - 3 \left[ \frac{M + (1/2)E^2}{M + (3/2)E^2} \right] \varepsilon^{(2)} - \frac{3}{2} \left[ \frac{M + E^2}{M + (3/2)E^2} \right] S \varepsilon^{(3)} + \varepsilon^{(4)}$$

$$> 0 \qquad < 0 \text{ if } \varepsilon^{(2)} > 0$$

The quantities  $\varepsilon^{(n)}$  represent the error made in  $\langle x^n \rangle / \sigma^n$  by ignoring the non-local dynamics.

From Diagonal Dominance, we expect that  $|\varepsilon^{(4)}| < |\varepsilon^{(3)}| < |\varepsilon^{(2)}|$  etc.

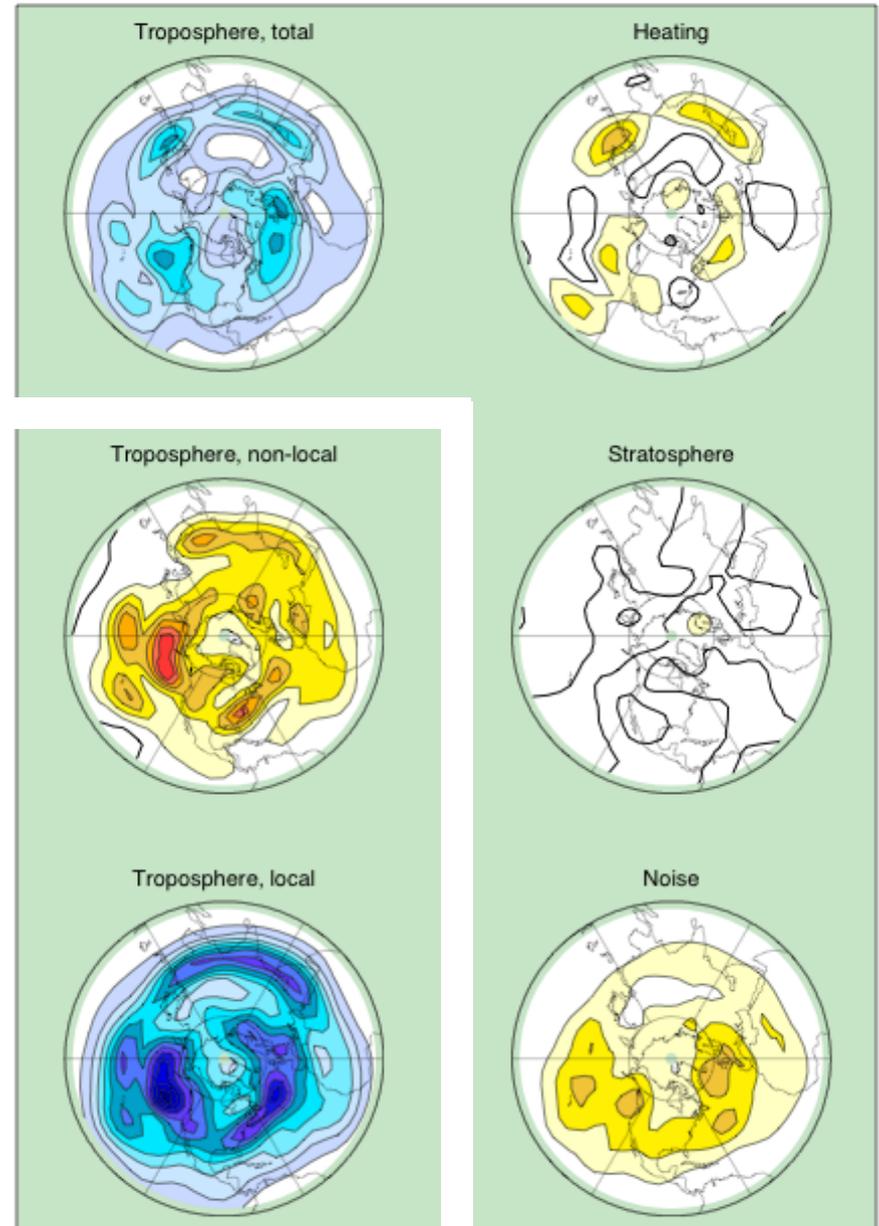
## Variance Budget of 250 mb Streamfunction in winter

Note the approximate  
balance between  
stochastic forcing and  
local damping.

*The non-local  
interactions increase the  
variance, everywhere.*

*Newman and  
Sardeshmukh  
(2008)*

250 hPa streamfunction variance budget



## Summary

1. Strong evidence for “coarse-grained” linear dynamics is provided by
  - ( a ) the observed decay of correlations with lag
  - ( b ) the success of linear forecast models, and
  - ( c ) the approximately linear system response to external forcing.
2. The simplest dynamical model with the above features is a linear model perturbed by **additive** Gaussian stochastic noise. **Such a model, however, cannot generate non-Gaussian statistics.**
3. A linear model with **a mix of multiplicative and additive noises** can generate non-Gaussian statistics; but not odd moments (such as skew) without external forcing; and therefore are not viable models of anomalies with zero mean.
4. **Linear models with correlated multiplicative and additive (“CAM”) noise** can generate both odd and even moments, and can also explain the remarkable observed quadratic K-S relationship between Kurtosis and Skew, as well as the Power-Law tails of the pdfs.