# Zonal Jets, Dipole EOFs, and Annular Modes

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- Obtained as EOFs, often treated as interchangeable
- Present study: how much of all of this can be understood from the kinematics of a fluctuating jet (without invoking complex dynamics)?



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$$C(x, x') = \mathsf{E}\left\{u(x, t)u(x', t)\right\} - \mathsf{E}\left\{u(x, t)\right\} \mathsf{E}\left\{u(x', t)\right\}$$



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EOFs orthogonal, PCs uncorrelated



## **Observed EOFs: Zonal Index and Annular Mode**



# **The Idealised Zonal Jet**

Assume eddy-driven midlatitude jet described by

$$u(x,t) = U(t)\mathcal{F}\left(\frac{x-x_c(t)}{\sigma(t)}\right)$$

$$\begin{split} U(t) &= U_0(1+l\xi(t)) & \text{jet strength} \\ \text{where} & x_c(t) = h\lambda(t) & \text{jet position} \\ & \sigma^{-1}(t) = 1+v\eta(t) & \text{jet width} \end{split}$$



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- Geopotential related to zonal wind through geostrophy:

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(where second term imposes mass conservation)

Define normalised basis functions  $F_j(x)$ :

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 so that

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$$\frac{d}{dx}F_j(x) = \frac{N_{j+1}}{N_j}F_{j+1}(x)$$

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and so

$$C(x, x') = U_0^2 h^2 N_1^2 F_1(x) F_1(x') + \frac{U_0^2}{2} N_1 N_2 h^3 s_\lambda [F_1(x) F_2(x') + F_2(x) F_1(x')] + \frac{U_0^2}{4} N_2^2 h^4(\kappa_\lambda + 3) F_2(x) F_2(x') + \dots$$

(where  $s_{\lambda}$ ,  $\kappa_{\lambda}$  skewness and kurtosis of  $\lambda$ )



• Writing EOF as  $E_u(x) = \alpha F_1(x) + \beta F_2(x)$  gives matrix equation (to  $O(h^4)$ ):

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- $s_{\lambda} \neq 0$  couples these basis functions in EOFs



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**Width Alone**: leading EOF pattern


## **Zonal Wind: Fluctuations in Single Variables**

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**Width Alone**: leading EOF pattern

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Leading PC time series couple position & strength fluctuations:

$$\alpha_u^{(1)}(t) \sim (U_0 + \xi(t))\lambda(t) + h.o.t.$$

For Gaussian jet with h = 0.3, l = 0.185





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Skewness in  $\lambda \Rightarrow$  EOFs asymmetric around jet axis; dipole still dominates even for strong skewness



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- For position, strength fluctuations of comparable width, coupling will be strong



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  - In fluctuations in  $\lambda$  not strongly skewed
  - position fluctuations not correlated with strength or width
  - position fluctuations stronger than width
- That these fairly general facts are characteristic of the tropospheric jet in models & observations ⇒ explanation for dipole as generic feature of zonal wind EOFs
- While dipole *arises* because of position fluctuations, the associated EOF mode bundles together variability in all jet degrees of freedom



Geopotential related to zonal wind through linear transformation

$$\Phi(x,t) = -\int_{x_1}^x f(x')u(x',t) \, dx' + \int_{x_1}^{x_2} \left( \int_{x_1}^x f(x')u(x')dx' \right) \mu(x) \, dx$$



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SO

$$C_{yy} = (L\mathcal{U})\Lambda(L\mathcal{U})^T$$

and EOFs of y only EOFs of x if rows of LU orthogonal

<u>CCCma</u> <u>Canadan Certre for Climate Modeling and Analysis</u> (which they won't be in general)



Can expand  $u'(x,t) = u(x,t) - \langle u(x) \rangle$  over EOF basis:

$$u'(x,t) = \sum_{j=1}^{J} \alpha_u^{(j)}(t) E_u^{(j)}(x)$$



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Mass conservation constraint influences non-orthogonality



## **Geopotential EOFs: Strength, Position, & Width**

Gaussian jet, neglecting sphericity of Earth ( $\mu(x) = f(x) = 1$ )





$$\Phi'(x,t) \simeq \alpha_u^{(1)}(t)G_1(x) + \alpha_u^{(2)}(t)G_2(x)$$



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$$\mathcal{I} = \int_{x_1}^{x_2} G_1(x) G_2(x) dx = -0.1$$









Leading EOF of  $\Phi(x, t)$  mixes EOFs of u(x, t)

<u>CCCma Canadam Centre for Climate Modeling and Analysis</u> <u>CCmac Centre canadien de la modélesation et de la modélesation et de la madeue de la modélesation et de la modélesation et de la madeue de la modélesation et de la m</u>

Annular mode structure requires both strength and position fluctuations and mixes leading two EOFs of zonal wind



Zonal Jets, Dipole EOFs, and Annular Modes - p. 18/20

# Geopotential: Strength, Position, & Width (Sphere)





# Conclusions

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